

Forcings and wave model results...







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Short course on ocean waves

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Outline

2. winds
4. currents
5. ice

Material : « Waves in geosciences » http://tinyurl.com/wavesgeo Stopa (Ocean Modelling 2018) Ardhuin et al. (Ocean Modelling 2014, JGR 2017)



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1



The action spectrum is defined by $N(k, \theta, \phi, \lambda, t) = E(k, \theta, \phi, \lambda, t) / \sigma$ And is the solution of a transport equation with **source term** :

$$\begin{split} \frac{\partial N}{\partial t} + \nabla_x \cdot \dot{\mathbf{x}} N + \frac{\partial}{\partial k} \dot{k} N + \frac{\partial}{\partial \theta} \dot{\theta} N = & \left[\frac{S}{\sigma} \right], \\ \dot{\mathbf{x}} = \mathbf{c}_g + \mathbf{U}, \\ \dot{k} = -\frac{\partial \sigma}{\partial d} \frac{\partial d}{\partial s} - \mathbf{k} \cdot \frac{\partial \mathbf{U}}{\partial s}, \\ \dot{\theta} = -\frac{1}{k} \left[\frac{\partial \sigma}{\partial d} \frac{\partial d}{\partial m} + \mathbf{k} \cdot \frac{\partial \mathbf{U}}{\partial m} \right], \end{split}$$



The source term represent many different effects...

- generation by the wind
- non-linear evolution
- dissipation due to breaking
- dissipation due to bottom friction
- scattering and reflection ... (this could have been in the LHS)

$$S = S_{ln} + S_{in} + S_{nl} + S_{ds} + S_{bot} + S_{db} + S_{tr} + S_{sc} + S_{ref} + S_{tr} + S_{sc} + S_{ref} + S_{tr} + S_{sc} + S_{ref} + S_{tr} + S_$$

Why separate different source terms ?

- different processes -> different parametrizations
- the source term ITSELF can be an interesting result ...



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Indeed, a source term S_a can be associated to

- a flux of energy phi_a = sum of (S_a df dth)
- (vector) flux of momentum tau_a,x = sum of (S_a/C cos(theta) df dth)

These fluxes are needed for coupling with ocean, ice, atmosphere





As a result the output of WW3 can be :

- full spectra (and associated source terms) + 2nd order option

- 1D spectral data : E(f), th1(f), p2l(f) ...

- « bulk » parameters (scalar / vector), with one value at each grid point. These are defined

- from the spectrum (e.g. Hs, Tm0,-1 ...)

- from the source terms (e.g. tauw, phioc ...)





Winds



Wind is usually quatified by the 10-minute average wind speed U₁₀ and direction at 10 m height

or by the wind stress vector... Wind speed measurements are more common... but not at 10m Wind is also influenced by waves :



Fig. 13. An LES snapshot of the instantaneous horizontal winds in a neutrally stratified boundary layer with mean 10-m wind speed of 5 m s⁻¹ and a wave age of c_p/U of 2.2. The color bar indicates the range of winds speeds 3.0 and 6.5 m s⁻¹. The wave height and wavelength are set to 1.6 and 100 m, respectively. Maximum wind speeds denoted by purple shading are found at a height of approximately 20 m.



Edson et al., J. Phys. Oceanogr. (2013)





Wind measurements can come from in situ or satellites... Wave models are usually driven by modeled winds

(not enough satellite data)



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For a long time the issue was resolution ...



Figure 8.3: Average wind and wave height in the Northern Hemisphere, Southern Hemisphere and Mediterranean sea, obtained in different runs of the ECMWF model in which the spatial resolution is changed. The resolutions of Gaussian grids T106, T213, T319, T511, T639, et T799 correspond to 188, 94, 63, 39, 31, and 25 km. Wind speed and wave height values have been divided by the values obtained for the coarsest of these grids. Taken from (Cavaleri and Bertotti, 2006, ©Elsevier).

Stopa (Ocean Modelling 2018)



But there are many other factors that improved the atmospheric models...



Janssen (Tech. Memo, ECMWF 2005)



But there are many other factors that improved the atmospheric models...



Figure 8.2: Root mean square error (RMSE) as a function of the forecast range, 0 corresponding to the analysis, in the operational systems ran at ECMWF. These modeled wind speed and wave heights are compared to buoy measurements over a year centered in winder (August to July). Only every other year is shown (Picture courtesy of J. Bidlot).

It is particularly important to get the high winds



Stopa (Ocean Modelling 2018)

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Given these uncertainties ... we need to tune the wave model to the different wind fields :

$$S_{\rm in}(k,\,\theta) = \frac{\rho_a}{\rho_w} \frac{\beta_{\rm max}}{\kappa^2} e^Z Z^4 \left(\frac{u_*}{C}\right)^2 \\ \times \max[\cos(\theta - \theta_u),\,0]^p \sigma F(k,\,\theta)$$







2 - Tidal currents can have a clear influence via refraction :



"Pierres Noires" buoy, WMO number 62069

Ardhuin et al. (J. Phys. Oceangr. 2012)



2 – Tidal currents can have a clear influence via refraction : Is it enough ?



In these two regions, it is possible that the current gradient is not well resolved in the tide model we use...

but how can we be sure that it is not just noise in the altimeter data ?



We expect western boundary currents to have strong impact on waves (Gulf Stream, Agulhas, Kuroshio, ...)



But we do not have measurements with high enough resolution :

- left : currents from satellite altimetry (Globcurrent)
- middle : measured wave heights and impact of these current in model
- right : modeled currents ... faster and sharper gradients.



But there are currents everywhere that will also impact the waves... Here is an example from Drake Passage (South of Chile) We use currents from the MITgcm IIc4320 simulation (see Rocha et al. 2017)



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Ardhuin et al. (J. Geophys. Res. 2017)



Summary

 Wind errors are the main sources of wave errors
There is no « ground truth » for wind speeds : we re-tune the wave model to different winds (or correct the winds) → Stopa (2018)

3) Currents are important too and can create large differences.

4) Currents can have effects on very small scale : wave height gradients are probably dominated by currents

5) Except for tidal currents, we have no good data for currents at scales < 200 km