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#### **Key Points:**

- 3-D approximate dispersion relation for surface waves on shear current whose magnitude and direction vary arbitrarily with depth
- The approximation is robust and accurate up to second-order with explicit estimates of error and criteria of applicability
- Compares favorably to approximation by Kirby and Chen (1989) whose properties and error estimate are derived for the first time

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### Approximate Dispersion Relations for Waves on Arbitrary Shear Flows

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**Abstract** An approximate dispersion relation is derived and presented for linear surface waves atop a shear current whose magnitude and direction can vary arbitrarily with depth. The approximation, derived to first order of deviation from potential flow, is shown to produce good approximations at all wavelengths for a wide range of naturally occuring shear flows as well as widely used model flows. The relation reduces in many cases to a 3-D generalization of the much used approximation by Skop (1987), developed further by Kirby and Chen (1989), but is shown to be more robust, succeeding in situations where the Kirby and Chen model fails. The two approximations incur the same numerical cost and difficulty. While the Kirby and Chen approximation is excellent for a wide range of currents, the exact criteria for its applicability have not been known. We explain the apparently serendipitous success of the latter and derive proper conditions of applicability for both approximate dispersion relations. Our new model has a greater range of applicability. A second order approximation is also derived. It greatly improves accuracy, which is shown to be important in difficult cases. It has an advantage over the corresponding second-order expression proposed by Kirby and Chen that its criterion of accuracy is explicitly known, which is not currently the case for the latter to our knowledge. Our second-order term is also arguably significantly simpler to implement, and more physically transparent, than its sibling due to Kirby and Chen.

**Plain Language Summary** In order to answer key questions such as how the ocean surface affects the climate, erodes the coastline and transports nutrients, we must understand how waves move. This is not so easy when depth varying currents are present, as they often are in coastal waters. We have developed a modeling tool for accurately predicting wave properties in such situations, ready for use, for example, in the complex oceanographic computer models. Our method is robust and works well in situations where the tool currently used will fail. In addition to predicting the speed of waves of different lengths and directions, it is important to know something about how accurate the prediction is, and as a worst case, whether it is reasonable at all. This has not been possible before, but we provide a way to answer both questions in a straightforward manner.

#### 1. Introduction

Knowledge of the properties of surface waves in the presence of currents is key to understanding, measuring and monitoring a range of processes in the oceans. Exchange of energy, mass and momentum between ocean and atmosphere is pivotal to predicting climate change, and local wave-current flow conditions can strongly affect the spread of nutrients as well as pollutants. Moreover, a broad range of fixed, floating or moored installations and vessels are affected by loads from waves and currents in combination; pertinent examples include floating oil booms which may fail under waves and currents, aquaculture farms in exposed locations, vessels, robots operating near the surface and even biolocomotion. Measuring currents from wave observations, e.g., using radar (Lund et al., 2015) is a favoured technique for obtaining flow field information toward such ends, but requires knowledge of how sheared currents affect wave dispersion. The effect of shear on waves is now included in widely employed oceanographic models such as Delft-3D (Elias et al., 2012) and ROMS, used for example in the coupled COAWST model (Kumar et al., 2011, 2012), and is recently employed for currents measurements using x–band radar observation of waves (Campana et al., 2017; Lund et al., 2015). A potentially important piece of progress toward incorporating arbitrary shear currents in oceanographic models was the recent derivation of an explicit wave action conservation equation

© 2017. American Geophysical Union. All Rights Reserved. on such currents by Quinn et al. (2017). They conclude that the neglect of shear in wave modeling in realistic ocean conditions can lead to significant errors.

In the coastal zone in particular, currents are often strongly sheared, and may change direction beneath the surface, for example when wind is blowing across a tidal current. In such conditions a fully 3-D approximate dispersion relation is required to correctly analyze the dispersion properties of surface waves, i.e., how the phase and group velocities of waves vary as a function of wavelength propagation direction and the shape and magnitude of the subsurface current. We present herein an approximation valid for an arbitrary horizontal velocity field  $\mathbf{U}(z)$  whose magnitude and direction may change as a function of *z*, providing good estimates of  $c(\mathbf{k})$  for all  $\mathbf{k}$  for typically occurring velocity profiles. This main result is found in equation (18).

Under many, but far from all, flow conditions of practical importance, the model coincides with a 3-D generalization of the much used approximation of Kirby and Chen (1989) to leading order in an expansion in terms of "deviation from potentiality," and involves the same calculational cost and complexity as the latter. However, several examples of realistic flows are presented where the Kirby and Chen model fails while our new one performs well or at least remains reasonable, demonstrating the improved robustness of our model. We also show that the criterion for our model to produce accurate approximations is less stringent than that of the Kirby and Chen approximation. Our derivations automatically produce a second-order accurate approximation—equation (19)—which is shown to be robust, and is arguably simpler and more transparent than the corresponding second-order approximation due to Kirby and Chen.

Although several approximations exist for the 2D and (with straightforward generalization) 3-D flow situation as reviewed below, some of which are excellent in most practical cases, there is a lack of understanding of their conditions of applicability and estimated error. We here rectify this situation by providing a thorough discussion of the accuracy and criteria of applicability of our approximation model as well as that of Kirby and Chen and a short-wavelength approximation due to Shrira (1993).

While work on waves on arbitrary shear currents is relatively sparse, a large body of literature exists on waves propagating on linearly varying currents in 2D. We make no attempt to review this very large literature, but mention but a few important papers. Key nonlinear treatments are those of Simmen and Saffman (1985) and Teles da Silva and Peregrine (1988). Of particular interest is the numerical scheme for fully nonlinear stationary and solitary waves on arbitrary 2D shear currents due to Dalrymple (1977), and followed by further numerical work on this question (Ko & Strauss, 2008a, 2008b; Moreira & Chacaltana, 2015; Nwogu, 2009). The present effort is restricted to linear waves; a comparison of the effect on wave speed of shear versus that of nonlinearity is an important topic for the future; refer to Swan and James (2000) for some results up to second order in wave steepness.

#### **1.1. Existing Approximation Methods**

Probably the most successful and widely used approximate dispersion relation is that due to Kirby and Chen (1989), which in turn was a direct generalization of previous models by Stewart and Joy (1974) and Skop (1987). Their approximation includes a second-order correction term, but we shall mostly be concerned with their first order correction, identical to that of Skop (1987). By the Kirby and Chen approximation the first order expression is meant, unless otherwise stated. We will work with a direct generalization of Kirby and Chen's model to 3D (3DKC). The success of Kirby and Chen's formula is likely in part to be due to the fact that, apparently serendipitously, the approximation it produces is often excellent even when the assumptions from which it was derived are strongly violated. In our analysis in the following we are able to explain this fortunate circumstance, which we believe has not previously been fully understood.

Among analytical dispersion relation approximations for arbitrary shear currents, that of Shrira (1993) is the only one to our knowledge explicitly derived for a fully 3-D configuration such as we consider herein, where the current may vary in both direction and magnitude with depth. Generalization of the Kirby and Chen approximation to 3D, however, is straightforward. *Shrira*'s relation is of limited use for oceanographic purposes because it is accurate only for very short waves. It is reviewed further in Appendix section A3.

An altogether different approach is to approximate the smooth velocity profile by a series of piecewise linear segments; we refer to this approach as the Piecewise-linear Approximation (PLA). This approach, which goes back a long time (e.g., Rayleigh, 1892; Taylor, 1955; Thompson, 1949), has the virtue that for a given wavevector **k** it converges as  $N^{-2}$  to the exact value of  $c(\mathbf{k})$  as the number of layers grows (Zhang, 2005), making it a useful comparison for the models developed here, and is preferable in cases where close control of errors, or particularly high accuracy, is required. The PLA was recently implemented by Smeltzer and Ellingsen (2017) in 3D, albeit with unidirectional shear current. The piecewise-linear approach is suitable when high and closely controlled accuracy for all wavelengths and directions is required. While numerically cheap, it remains more cumbersome to implement than analytical approximations due to the need to eliminate additional, spurious solutions (Smeltzer & Ellingsen, 2017).

Some further approaches to approximating wave dispersion on curved velocity profiles should be mentioned, with no ambition of completeness. Zhang (2005) presents an analytical approximation which amounts to a matching of the short-wavelength model of *Shrira* to the 3DKC model at longer wavelengths. Solution schemes have also been developed for specific velocity profiles; exponential current profiles modeling wind-driven currents were studied by Abdullah (1949) and the 1/7 law profile by Hunt (1955) and Fenton (1973). Several examples are worked out with a method developed by Karageorgis (2012). The case of shallow waters compared to wavelength was considered by Burns (1953). Numerical approaches, using shooting methods, have been used e.g., by Fenton (1973) and Dong and Kirby (2012).

#### 1.2. Approach

The widely used approximation by Kirby and Chen (1989), like its predecessors by Stewart and Joy (1974) and Skop (1987), all proceed by performing a formal expansion in orders of a dimensionless parameter which is assumed small *a priori*. Our present approach is somewhat different, and is inspired by that of Shrira (1993). We adopt a 'near–potentiality' assumption, that the wave–induced fluid motion does not differ greatly from that due to a flow with linearly varying velocity profile. The formal procedure is to isolate the effect of the curvature of the velocity profile in a single term in an implicit non–closed dispersion relation, and combine it with an iterative solution to the Rayleigh equation by a method of dominant balances assuming the term resulting from curvature to be dominated. Like Shrira (1993) a 'true' corresponding small parameter comes out as a *result* rather than an initial assumption, and may be interpreted as a suitably depth–averaged measure of the shear-profile's curvature.

#### 1.3. Outline

In the next section our approximate dispersion relation, equation (18), is presented and analyzed. We analyze the question of the range of applicability of our perturbation scheme, providing order of magnitude estimates for the error. The relation to the 3DKC is discussed, including the conditions under which the two coincide to leading order. Corresponding approximations for group velocity are also presented. A brief discussion of critical layers is given in section 3. In section 4 we test our model for a range of realistic shear profiles, both unidirectional and varying in direction as a function of depth, including particular cases where the 3DKC breaks down. Summary and conclusions are found in section 5, and some further information and details may be found in Appendix A.

#### 2. An Approximate Relation for 3-D Wave-Shear Current Dispersion

We consider the system shown in Figure 1. A horizontal current  $\mathbf{U}(z)$  whose direction and magnitude may vary with depth, running over a flat sea bed of depth h and with a free surface which is at z = 0 when undisturbed. Consider a plane wave with wave vector  $\mathbf{k} = [k_x, k_y]$  propagating atop the current. The wave creates a disturbance of velocity and pressure fields. We work within linear wave theory, hence all equations of motion and boundary conditions are linearized with respect to perturbations due to the wave motion. All perturbations are understood to be proportional to exp  $[i\mathbf{k} \cdot \mathbf{x} - ikc(\mathbf{k})t]$ ; here  $\mathbf{x} = [x, y]$  is the position in the horizontal plane, c is the phase velocity in direction  $\mathbf{k}$ ,  $k = |\mathbf{k}|$ , and t is time. We neglect surface tension and viscosity, so the flow is governed by the Euler equation, which reduces to the following boundary value problem for the vertical velocity w(z) only (e.g., Shrira, 1993),

$$w''(z) - k^2 w(z) = \frac{\mathbf{k} \cdot \mathbf{U}''(z)}{\mathbf{k} \cdot \mathbf{U}(z) - kc} w(z);$$
(1a)

$$\mathbf{k} \cdot \mathbf{U}_0 - kc)^2 w'(0) - [\mathbf{k} \cdot \mathbf{U}_0'(\mathbf{k} \cdot \mathbf{U}_0 - kc) + gk^2] w(0) = 0;$$
(1b)

$$w(-h)=0.$$
 (1c)

Equation (1a) is called the Rayleigh equation (or inviscid Orr–Sommerfeld equation), and equations (1b) and (1c) are the appropriate boundary conditions at the free surface and bottom, respectively. We defined  $U_0 =$ 



Figure 1. The geometry: a plane wave of wave vector **k** travels atop a horizontal shear current whose magnitude and direction may vary with depth.

 $\mathbf{U}(0)$  and  $\mathbf{U}'_0 = \mathbf{U}'(0)$ , and g is gravitational acceleration. For future reference we define velocities relative to the surface

$$\tilde{c} = c - \mathbf{k} \cdot \mathbf{U}_0 / k; \ \Delta \mathbf{U}(z) = \mathbf{U}(z) - \mathbf{U}_0.$$
<sup>(2)</sup>

Here  $\tilde{c}$  is the intrinsic phase velocity.

Our task in this section will be to derive, discuss and test the approximate dispersion relation  $c(\mathbf{k}) \approx c_{\approx}(\mathbf{k})$  which is a main result and is given in equation (18).

Inspired by Shrira (1993) we make a 'near–potentiality' assumption which entails that the effect of the current on the dispersion relation differs only moderately from the explicitly solvable case of a linearly varying  $\mathbf{U}(z)$ . On the other hand, the shear itself need not be small. (Note that the term 'near–potentiality' is a slight misnomer since a wave propagating at an oblique angle with such a current in fact has vorticity and is not as such 'potential'; see Ellingsen (2016). Potential theory can be employed to waves on linearly depth– dependent currents only in strictly 2-dimensional flow (e.g., Ellingsen and Brevik, 2014).)

#### 2.1. Recasting the Equations of Motion

We proceed in two steps. First, the boundary value problem (1) is written in an alternative way; combining the boundary conditions and Rayleigh equation we find an implicit dispersion relation (8) with both  $\tilde{c}$  and w unknown, where the effect of the depth–averaged shear is separated from that due to higher derivatives of  $\mathbf{U}(z)$ . Approximations of w are inserted into this equation, found by iterative solutions of (1a) presuming the right–hand side to be a dominated term. In section 2.2 we will term the approximate dispersion relation accounting only for depth–averaged shear the 'first order' approximation, and the improved approximation from including the term in (8) due to curvature as well as the second–level approximation for w, we term the 'second order' approximation of  $\tilde{c}$ .

A fruitful starting point is to frame the boundary value problem (1) in the form,

$$[1+I(\tilde{c})]k^2\tilde{c}^2 + \tilde{c}\mathbf{k}\cdot\mathbf{U}_0'\tanh kh - k^2c_0^2 = 0$$
(3)

with

$$I(\tilde{c}) = \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z)w(z)\sinh k(z+h)}{kw(0)[\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}]\cosh kh} dz$$
(4)

and  $c_0 = \sqrt{(g/k) \tanh kh}$ . This is obtained by multiplying (1a) by  $\sinh k(z+h)$ , integrating with respect to *z*, and inserting boundary conditions (1b) and (1c). This very useful relation informs most of our analysis in this paper. In itself the relation (3) is not closed since both *w* and  $\tilde{c}$  are unknown.

Equation (3) appears to isolate the effect of the curvature  $\mathbf{U}''$  in the term *I*, but this is deceptive: there is still an explicit reference to the surface shear  $\mathbf{U}'_0$  which may also be re-written as a depth–integrated curvature. The appearance of  $\mathbf{U}'_0$  is moreover undesirable: long waves in particular must be expected to be influenced by the suitably depth–averaged shear, wheras the value at the surface is comparatively irrelevant.

To transform (3) into the desired form we write the integral  $I(\tilde{c})$  as

$$I(\tilde{c}) \equiv I_0 + \frac{c_0^2}{\tilde{c}^2} \varepsilon I(\tilde{c}) = c_0 \tilde{c} \left( K - \frac{\mathbf{k} \cdot \mathbf{U}_0'}{k^2 c_0} \tanh kh \right) + \frac{c_0^2}{\tilde{c}^2} \varepsilon I(\tilde{c}), \tag{5}$$

with

$$\varepsilon I(\tilde{c}) = \frac{\tilde{c}}{c_0^2} \int_{-h}^{0} \frac{(\mathbf{k} \cdot \mathbf{U}'')(\mathbf{k} \cdot \Delta \mathbf{U})w(z)\sinh k(z+h)}{k^2(\mathbf{k} \cdot \Delta \mathbf{U} - k\tilde{c})w(0)\cosh kh} dz;$$
(6)

$$\mathcal{K} = \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}'}{k^2 c_0 \cosh kh} \frac{\mathrm{d}}{\mathrm{d}\zeta} \left[ \frac{w(\zeta)}{w(0)} \sinh k(\zeta + h) \right]_{\zeta = z} \mathrm{d}z.$$
(7)

Here  $I_0$  is obtained from I by setting  $\Delta \mathbf{U} = 0$ , and the last form is obtained by partial integration of  $I_0$ . Inserting into (3) gives

$$\tilde{c}^{2} + c_{0}\tilde{c}K + c_{0}^{2}\varepsilon I(\tilde{c}) - c_{0}^{2} = 0$$
(8)

which no longer contains an explicit reference to the surface. Instead the effect of the shear is contained in K, and the effect of the curvature shared between K and  $\varepsilon I$ .

Equation (8) is exact within linear wave theory, and we now begin making approximations. To wit we wish to isolate the part of *K* which depends on the shear but not the curvature, plus a curvature–related correction. We adopt an iterative approach rather than a formal expansion in an explicit parameter as done by Kirby and Chen (1989) and predecessors. Noting that the right–hand side of (1a) is proportional to the curvature  $\mathbf{U}''(z)$ , we presume its influence on *w* and  $\tilde{c}$  be small, and that successively better approximations are obtained by iterative solutions by method of dominant balance (cf. e.g., Bender & Orszag, 1978). This is our understanding of the 'near–potentiality' assumption. We let  $w(z) = w^{(0)}(z) + \varepsilon w^{(1)}(z) + \cdots$  and verify that

$$w^{(0)}(z) = w^{(0)}(0) \frac{\sinh k(z+h)}{\sinh kh},$$
(9a)

$$\varepsilon w^{(1)}(z) = \frac{w^{(0)}(0)}{k} \int_{-h}^{z} \frac{\mathbf{k} \cdot \mathbf{U}''(\zeta)}{\mathbf{k} \cdot \Delta \mathbf{U}(\zeta) - k\tilde{c}} \frac{\sinh k(\zeta + h)\sinh k(z - \zeta)}{\sinh kh} d\zeta.$$
(9b)

Inserting the leading approximation of w(z) we obtain  $\varepsilon I = \varepsilon \Omega_I + \cdots$  with

$$\varepsilon \Omega_{l}(\tilde{c}) = \frac{2\tilde{c}}{k^{2}c_{0}^{2}} \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z)\mathbf{k} \cdot \Delta \mathbf{U}(z)\sinh^{2}k(z+h)}{[\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}]\sinh 2kh} dz.$$
(10)

Now inserting the expansion for w(z) into (7) gives, after some algebra,  $K = 2\delta + (c_0/\tilde{c})(\epsilon\Omega_{K1}\delta + \epsilon\Omega_{K2})$  with  $\delta$  defined in (12), and

$$\epsilon \Omega_{K1}(\tilde{c}) = \frac{2\tilde{c}}{kc_0} \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z)}{\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}} \frac{\sinh k(z+h)\sinh kz}{\sinh kh} dz, \qquad (11a)$$

$$\epsilon \Omega_{K2}(\tilde{c}) = \frac{2\tilde{c}}{k^2 c_0^2} \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z)}{[\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}]} \frac{\sinh k(z+h)}{\sinh 2kh} \\ \times \left[ \int_{z}^{0} \mathbf{k} \cdot \mathbf{U}'(\zeta) \sinh k(2\zeta+h-z) d\zeta \right] dz. \qquad (11b)$$

We have introduced a key parameter in our theory,

$$\delta(\mathbf{k}) = \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}'(z) \sinh 2k(z+h)}{kc_0 \sinh 2kh} dz,$$
(12)

which is a dimensionless depth-averaged shear, to which we shall refer frequently.

We thus obtain the approximation for (8) on the desired form

$$\tilde{c}^2 + 2c_0\tilde{c}\delta - c_0^2 + c_0^2\Delta(\tilde{c}) + \dots = 0$$
(13)

with  $\Delta(\tilde{c}) = \epsilon \Omega_{l}(\tilde{c}) + \epsilon \Omega_{K1}(\tilde{c}) \delta + \epsilon \Omega_{K2}(\tilde{c})$ , which may be simplified with some tedious manipulation (see Appendix section A2) to the form

$$\Delta(\tilde{c}) = -\frac{2\tilde{c}}{kc_0^2} \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z)}{\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}} \frac{\sinh k(z+h)\sinh kz}{\sinh kh} \left[\tilde{U} - \tilde{u}(z)\right] dz, \tag{14a}$$

$$\tilde{u}(z) = -\frac{\sinh kh}{\sinh kz} \int_{z}^{0} \frac{2\mathbf{k} \cdot \mathbf{U}(\zeta) \cosh k(2\zeta + h - z)}{\sinh 2kh} d\zeta.$$
(14b)

The weighted depth-averaged velocity, as introduced by Skop (1987) and used extensively by Kirby and Chen (1989), is

$$\tilde{U}(\mathbf{k}) = \frac{\mathbf{k} \cdot \mathbf{U}_0}{k} - c_0 \delta(\mathbf{k}) = \int_{-h}^0 \frac{2\mathbf{k} \cdot \mathbf{U}(z) \cosh 2k(z+h)}{\sinh 2kh} dz$$
(15)

(the middle form is found by partial integration). Inspection reveals that  $\tilde{u}(z)$  is a possible generalization of the depth averaged velocity  $\tilde{U}$  when averaging is only carried out down to a depth z rather than the full depth. In particular,  $\tilde{u}(-h) = \tilde{U}$  and  $\lim_{z \to 0} \tilde{u}(z) = \mathbf{k} \cdot \mathbf{U}_0/k$ .

The effect of shear is now contained in the dimensionless mean shear  $\delta$ , and the effect of curvature of  $\mathbf{U}(z)$  has successfully isolated in the dimensionless quantity  $\Delta(\tilde{c})$ . In practice,  $\Delta(\tilde{c})$  is calculated using an approximate value of  $\tilde{c}$ , as will be detailed below. If  $\mathbf{k} \cdot \Delta \mathbf{U} = k\tilde{c}$  somewhere in the integration range, the outer integral in (14) has a pole and the principal value should be taken (Shrira, 1993). Deep water limits of  $\delta$  and  $\Delta$  are quoted in Appendix section A1.

Equation (13) now makes explicit what the true small parameter is, namely  $\Delta(\tilde{c})$ . Note that, like Shrira (1993), we find that the effect of the curvature, corresponding to the right–hand side of (1a), only need be small in a depth–integrated sense.

#### 2.2. New Approximation for $\tilde{c}$

Solving (13) with respect to  $\tilde{c}$  gives

$$\tilde{c} \approx c_0 \left[ \sqrt{1 + \delta^2 - \Delta} - \delta \right]. \tag{16}$$

In accordance with the 'near-potentiality' assumption we presume  $\Delta \ll 1$  while making no assumptions about the strength of the (non-dimensional) averaged shear. This gives a new approximation

$$c \approx c_{\approx} + \tilde{c}_{\approx,2nd}$$
 (17)

with the leading order approximation

$$c_{\approx}(\mathbf{k}) = \frac{\mathbf{k} \cdot \mathbf{U}_0}{k} + c_0 \left(\sqrt{1 + \delta^2} - \delta\right), \tag{18}$$

which is our main result, and a second order correction

$$\tilde{c}_{\approx,2nd} = -\frac{c_0 \Delta(\tilde{c}_{\approx})}{2\sqrt{1+\delta^2}}.$$
(19)

To calculate  $\Delta$  in practice, the first order estimate is inserted for  $\tilde{c}$ , as indicated. We test the new approximation  $c_{\approx}$  in a range of cases in section 4.

A few remarks about the approximation (18) are warranted. The first of these is that when  $\mathbf{U}(z)$  is a linear function of z, i.e.,  $\mathbf{U}(z) = \mathbf{U}_0 + \mathbf{U}'_0 z$ , equation (18) gives the *exact* answer which is well known to be  $\tilde{c} = c_s$  (e.g., Ellingsen, 2014), with

$$c_{s}(\mathbf{k}) = -\frac{\mathbf{k} \cdot \mathbf{U}_{0}'}{2k^{2}} \tanh kh + \sqrt{c_{0}^{2} + \left(\frac{\mathbf{k} \cdot \mathbf{U}_{0}'}{2k^{2}} \tanh kh\right)^{2}}.$$
(20)

(This relation may have been given first by Craik (1968)). The majority of the analytical literature on waves on shear flow concerns this type of flow, for which reason it is an important benchmark.

We second remark that the above derivations were performed for an arbitrary  $\mathbf{k}$ . No assumptions were made which might restrict the theory to a particular range of wavelengths.

Third one may note that the widely used (3D) Kirby and Chen approximation (22), which we will review in a moment, is conventionally implemented by calculating  $\tilde{U}$  rather than  $\delta$ . If this is preferred,  $c_{\approx}$  is expressed in terms of  $\tilde{U}$  using (15):

$$\boldsymbol{c}_{\approx}(\mathbf{k}) = \sqrt{\boldsymbol{c}_{0}^{2} + (\tilde{\boldsymbol{U}} - \mathbf{k} \cdot \mathbf{U}_{0}/\boldsymbol{k})^{2} + \tilde{\boldsymbol{U}}},$$
(21)

Numerical implementation of (18) and the 3DKC thus involves essentially identical complexity and calculational effort.

Finally, we note that  $\tilde{c}_{\approx}$  is positive for all values of  $\delta$  as physically it must be.

#### 2.3. Comparison With the Approximation of Skop/Kirby and Chen

We observe that in the special case  $\delta \ll 1$  equation (18) coincides, modulo a term of order  $\delta^2$ , with the 3DKC:

$$c_{\mathsf{KC}}(\mathbf{k}) = \frac{\mathbf{k} \cdot \mathbf{U}_0}{k} + c_0(1-\delta) = c_0 + \tilde{U}$$
(22)

The weighted depth averaged velocity  $\tilde{U}$  was defined in (15). This first order expression (in orders of  $kc/\mathbf{k} \cdot \mathbf{U}$ ) was derived by Skop (1987), and extended to next order by Kirby and Chen (1989). In contrast with approximation (18), the 3DKC can predict nonsensical, negative values for  $\tilde{c}$  when  $\delta > 1$ , corresponding to strongly sheared flows.

It is clear from this, and further discussed below, that the 3DKC works well when  $\delta$  is small compared to unity. In certain important cases  $\delta(\mathbf{k})$  may *not* be small, in which case (18) is superior to (22) except in special cases where cancellations occur, as will be explained in section 2.4. Indeed, should  $\delta(\mathbf{k}) > 1$ , the 3DKC predicts  $\tilde{c} < 0$  which is physically unacceptable.

From (16) a second order term may also be derived, equal to  $c_0(\delta^2 - \Delta)/2$ . An "extended KCA" is then found by adding this term to to (22), an alternative to the second order expression proposed by Kirby and Chen (1989). Unlike (22) the "extended KCA" is always positive, and is second order accurate when  $\delta$  is small, but (19) is typically superior when  $\delta$  approaches unity. We have not studied this approximation in detail.

Compared to the second-order approximation of Kirby and Chen (1989), (19) has the clear advantage that it makes no assumptions about the length or velocity of waves, nor the strength of the shear, and requires the same criterion as the leading approximation to hold, given in (24) below, making it well controlled. Our examples indicate that the second-order expression by Kirby and Chen (1989) is generally robust and accurate, but we have not succeeded in identifying the criteria for this to hold true. (19) is also arguably a simpler and more transparent expression than the second-order correction derived by Kirby and Chen, and our own experience is that it is significantly easier to implement, while admitting that this is to an extent a point of preference.

#### 2.4. Applicability Criteria and Error Estimates

Our goal in this section is two-fold. First we derive the pertinent applicability criteria for the new approximation (18) as well as the 3DKC (22); these are found in equation (24) through (26). We use these results to explain the surprising success of the 3DKC, not previously understood to our knowledge — Indeed it seems to us that the sundry applicability ranges at play have been a source of some puzzlement, evident, for example, in the discussions of Swan and James (2000). The applicability of the short-wave approximation due to Shrira (1993) is discussed in Appendix section A3.

We shall see in the following that inclusion of the second order term in (17) strongly increases accuracy for all wavelengths in the examples tested. With the exception of two extremely strongly sheared flows considered in section 4.4 (where, we shall see, applicability criteria are not satisfied), the estimate (16) is accurate

to a percent or less. The criterion that  $\tilde{c}_{\approx}$  from (18) is a good approximation we then take to be that  $(\tilde{c} - \tilde{c}_{\approx}) / \tilde{c}_{\approx} \ll 1$  where we use (16) for the "exact" value. For approximation (18) this means

$$|\Delta| \ll 2 \left| \delta \sqrt{1 + \delta^2 - 1 - \delta^2} \right|. \tag{23}$$

The right hand side is strictly greater than 1 for all  $\delta$ . A sufficient and far simpler criterion for this to hold is that

$$|\Delta(\tilde{c})| \ll 1. \tag{24}$$

Since  $|\Delta|$  is a measure of the effect of the right–hand side of the Rayleigh equation (1a) on  $\tilde{c}$ , i.e., the effect of the curvature of the velocity profile, criterion (24) is a formal 'near–potentiality' criterion.

In contrast, a sufficient criterion of applicability of the 3DKC (22) based on the same argument is that

$$\left. \frac{\delta^2 - \Delta}{2(1 - \delta)} \right| \ll 1 \tag{25}$$

assuming  $\delta < 1$ .

A sufficient criterion for this to hold is the *double* criterion

$$\delta^2 \sim |\Delta(\tilde{c})| \ll 1. \tag{26}$$

Particularly, in cases where  $|\Delta| \ll 1$  is satisfied but  $\delta$  is not small, the new approximation (18) is most often superior. However, (25) shows that situations exist where the 3DKC happens to be accurate even when  $\delta$  is not so small, because another sufficient criterion implying (25) is that

$$|\delta^2 - \Delta| \ll 1,\tag{27}$$

provided  $2(1-\delta) \sim O(1)$ . In such cases a fortunate cancellation occurs in the next order correction to (22), rendering the first order approximation accurate even when  $\delta^2$  and  $\Delta$  are not so small respectively, as we shall see in numerical examples in section 4.

#### 2.4.1. The Broad Applicability of 3DKC Explained

Remarkably, the 3DKC (22) can be derived by *either* assuming  $kc \gg |\mathbf{k} \cdot \mathbf{U}|$  (Kirby & Chen, 1989; Stewart & Joy, 1974) (The assumption  $kc \gg |\mathbf{k} \cdot \mathbf{U}|$  is not in general satisfactory since  $\mathbf{U}$  and c (unlike  $\Delta \mathbf{U}$  and  $\tilde{c}$ ) depend on the choice of reference system. In many cases this is easily rectified in a way amounting to the same analysis.) *or* by taking the in some sense opposite limit  $k \to \infty$ . This surprising coincidence was noted by Skop (1987), but seems not to have been realized by either Kirby and Chen or Stewart and Joy. We are now in a position to provide an explanation.

Inspecting the definition (12), and provided 'near potentiality' (24) is satisfied, the criterion  $\delta^2 \ll 1$  from (26) implies that the 3DKC is sure to be excellent in three different limiting cases where  $\delta$  vanishes:

- 1. Weak current,  $kc_0/\mathbf{k} \cdot \mathbf{U} \rightarrow \infty$ ;
- 2. When the shear is weak,  $\mathbf{k} \cdot \mathbf{U}'(z)/k^2 c_0 \rightarrow 0$ ;
- 3. When wavelength is short,  $k \to \infty$ .

That  $\delta \to 0$  in the first case may be seen from (22), whereby  $\delta = \mathbf{k} \cdot \mathbf{U}_0/kc_0 - \hat{U}/c_0$ . The third follows from the large *k* asymptotic  $\delta \sim \mathbf{k} \cdot \mathbf{U}_0'/(2k^2c_0) \sim \mathbf{k} \cdot \mathbf{U}_0'/(2g^{\frac{1}{2}}k^{\frac{3}{2}})$  which tends to zero in this limit.

Kirby and Chen (1989) derive (the 2D case of) equation (22) based on, essentially, the first of these cases, but conjecture that the true condition might be weak shear. Skop (1987) remarks that essentially the same approximation may be derived assuming short waves instead. These are exactly the three regimes listed. **2.4.2. Error Estimates** 

Noting that the second order approximation constitutes a significant improvement in accuracy for  $\tilde{c}$  in all examples considered, we propose that its magnitude relative to the first-order approximation of  $\tilde{c}$  provides a rough estimate of the error of using (18) and (22), respectively,

$$\operatorname{err}_{\approx} \sim \frac{c_{0}|\Delta(c_{\approx})|}{2\tilde{c}_{\approx}\sqrt{1+\delta^{2}}}, \quad \operatorname{err}_{\mathsf{KC}} \sim \frac{|\delta^{2} - \Delta(\tilde{c}_{\mathsf{KC}})|}{2(1-\delta)}.$$
(28)

In general,  $err_{\approx}$  is of order  $\Delta$ , while  $err_{KC}$  is of order max ( $\delta^2$ ,  $\Delta$ ). However, as mentioned the 3DKC can conspire to be smaller than this due to partial cancellation between  $\delta$  and  $\Delta/\delta$  ( $err_{KC}$  is only valid for  $\delta$  smaller than, and not too close to, 1.)

#### 2.5. Group Velocity

The approximate group velocity according to the approximation (18) is found as

$$c_{g,\approx}(\mathbf{k}) = \frac{\mathrm{d}}{\mathrm{d}k} k c_{\approx}(\mathbf{k}) = \frac{\mathbf{k} \cdot \mathbf{U}_0}{k} - C_L + (1 + \delta^2)^{-1/2} (c_{g0} + C_L \delta),$$
(29)

where  $C_L = \frac{d}{dk}(kc_0\delta)$ , and  $c_{g0} = \frac{1}{2}c_0(1+2khcsch2kh)$  is the group velocity with no current. Explicitly

$$C_L = (1 - 2kh \coth 2kh)c_0 \delta + 2 \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}'(z)(z+h) \cosh 2k(z+h)}{\sinh 2kh} dz.$$
(30)

For reference, the group velocity predicted with 3DKC (22) is  $c_{g,KC} = \mathbf{k} \cdot \mathbf{U}_0 / k + c_{g0} - C_L$ . A comprehensive discussion of the group velocity using the 3DKC was recently given by Banihashemi et al. (2017).

#### 3. Application to Critical Layers

Especially in the presence of strong surface shear, critical layers could occur near enough to the surface to affect wave properties. Critical layers occur when a critical depth  $z_c$  exist so that  $\mathbf{k} \cdot \mathbf{U}(z_c) = kc$ . In the presence of a critical layer the eigenvalues for the phase velocity c acquire an imaginary part whose sign determines the stability of the flow. A thorough background is provided by Drazin and Reid (1981). A fully realistic treatment of this problem requires inclusion of viscous effects (Velthuizen and van Wijngaarden, 1969), which is beyond the scope of this study. The inviscid problem is nevertheless an informative model and much of the literature concerns stability of the Rayleigh equation (1a) rather than the full Orr–Sommerfeld problem (Morland et al., 1991; Young & Wolfe, 2014).

We show in Appendix section A4 how equation (3) allows us to generalize the approximate treatment by Shrira (1993) in a simple way. We conclude, like Shrira (1993), that instability is predicted when  $\mathbf{k} \cdot \mathbf{U}''(z_c) < 0$ , but derived here under far less restrictive assumptions. The explicit approximation for Im(c) is given in Appendix section A4. The prediction is made by linearizing with respect to  $\Delta/\delta$ , and is valid only up to linear order in this parameter.

#### 4. Numerical Results

In this section we test our approximation (18) and (19) for different shear flows, with special emphasis on error estimates and the smallness parameters considered in section 2.4. For all flows we have also calculated  $\tilde{c}(\mathbf{k})$  with high accuracy using the piecewise-linear approximation (PLA) detailed in Smeltzer and Ellingsen (2017), allowing us to accurately compare with the "exact" answer for all flows considered. The PLA calculations all have relative accuracy better than  $10^{-4}$  for intrinsic phase velocities.

#### 4.1. Typical Wind-Driven Flows

Flows with strong shear near the surface will affect the dispersion the most. A typical example is winddriven flow. Examples of wind driven flow profiles were collected and analyzed by Swan and James (2000) (see further references therein); we use the velocity profiles from their Figure 2, assuming waves propagating downstream. Results are shown in Figure 2.

For all of these flows the new approximation (18) and the 3DKC (22) are essentially equally good, which accords with our analysis in section 2.4 since  $\delta$  remains significantly less than 1 for all *k*. Exactly which of the two happens to be closest to the exact solution for some *k* depends on the exact shape of U(z), and we do not see a way to predict this short of a full calculation. It is notable that the 3DKC works well even for  $\delta$  as large as 0.35, which is due to the partial cancellation between  $\delta^2$  and  $\Delta$ , criterion (25).

Both our second-order approximation (19) and the higher order expression due to Kirby and Chen (1989) improve accuracy, and are essentially equally good for these moderately sheared flows.

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**Figure 2.** Approximate dispersion relations applied to different wind-drift shear currents (Swan & James, 2000). (a, b, c) Results for the shear profiles are found in their respective columns. (d, e, f) Calculated estimates of intrinsic velocity  $\tilde{c}$  relative to the "exact" value calculated with the piecewise–linear method (PLA). Results are calculated for the 1st and second-order approximations found herein in equations (18) and (19), respectively, and those due to Kirby and Chen (1989). (g, h, i) The parameters relevant to applicability of the approximations.  $\Delta(\tilde{c})$  were calculated from the PLA value of  $\tilde{c}$ .

We have performed the same calculation for a shear–assisted wave, equivalent to letting  $\mathbf{U} \rightarrow -\mathbf{U}$  in Figure 2. The same conclusions hold, both first order expressions are accurate to better than 3%, and both second-order forms are much better than this.

#### 4.2. Comparison With Exact Cases

We go on to test the different approximations for a particular class of shear currents analyzed by Peregrine (1976),  $\mathbf{U}(z) = U(z)\mathbf{e}_x$  with

$$U(z) = U_0 \cosh \kappa z + U'_0 \kappa^{-1} \sinh \kappa z.$$
(31)



**Figure 3.** Approximate dispersion relations used for waves with exact phase velocity c = 0 on profile (31), for different propagation directions  $\theta$ . In all plots,  $U_0/\sqrt{gh}=0.45$ .  $hU'_0/U_0=$  (a-d) 0 and (e-h) 4;  $\kappa$  is determined from the condition U(-h)=0. (a, e) Velocity profiles for each row. (b, f) Phase velocities from different approximations; see Figure 2 for abbreviations. (c, g) Comparison of the different smallness parameters considered in section 2.4. Wave number  $k(\kappa, U_0, U'_0)$  is calculated from equation (4.21) of Peregrine (1976), and plotted in Figures 3d and 3h.

For the specific case c = 0, the Rayleigh equation (1a) can be solved exactly giving  $w(z)=w(0)\sinh K(z+h)/\sinh Kh$  with  $K=\sqrt{k^2+\kappa^2}$ . The wave number k is a function of the three parameters  $\kappa$ ,  $U_0$  and  $U'_0$  and solves an implicit dispersion relation. See (Peregrine, 1976, pp. 79–82) for full details. We choose U(-h)=0, which fixes  $\kappa$  implicitly.

Results are shown in Figure 3, where two different profiles of type (31) are analyzed with  $\theta$  being the angle between **k** and **U**; velocity profiles are shown in panels (a) and (e), and each row presents calculations pertaining to their respective profiles. The value of *kh* which corresponds to c = 0 is shown in panels (d) and (h), showing that the wavelengths become very short as  $\theta$  approaches  $\pi/2$ , whereas the waves near  $\theta = -\pi$  the waves have wavelength comparable to *h*, and are thus affected by the flow within a significant portion of the water column. The small values of  $\delta$  for the profile in Figure 3a mean the two predictions are almost identical and both predict  $c \sim 10^{-3} U_0$ . For the profile in Figure 3e with strong surface shear,  $\delta$  is higher, yet there is no significant difference in prediction accuracy between the two models, which can be explained by the partial cancellation between  $\Delta$  and  $\delta^2$  in err<sub>KC</sub> in equation (28).

#### 4.3. Fully 3-Dimensonal Flows

In realistic settings, the direction of  $\mathbf{U}(z)$  may vary with depth. An example could be a wind-driven flow across a tidal current. As a model flow we use an exponential/trigonometric profile

$$\mathbf{U}(z) = U_0 \sinh \alpha (z+h) [\cos (\kappa z + \phi) \mathbf{e}_x + \sin (\kappa z + \phi) \mathbf{e}_y]$$
(32)

where  $\alpha$  and  $\kappa$  are parameters for the vertical and horizontal depth variation. Results are shown in Figure 4 for different propagation angles  $\theta$  and three different values of the parameter  $\kappa h$  ranging from  $\kappa h=0$  (no directional variation) to  $\kappa h=4$  (strongest directional variation) as shown. We use kh = 1 everywhere, and  $\theta$  is the angle between **k** and **U**(0).

There are several interesting phenomena to note. Panel (a) shows significant discrepancy between our model (18) and the 3DKC, increasing with stronger vertical shear, in the vicinity of  $\theta = 0$ . We do not have "exact" calculations to compare with for the direction-changing profile, yet the previous error and applicability analysis gives an explanation of discrepancies when the key parameters  $\delta$  and  $\Delta$  are also considered,



**Figure 4.** Demonstration of approximate dispersion relation for the direction-changing flow (32) with  $U_0/\sqrt{gh}=0.5$ ,  $\alpha h=1$ ,  $\phi=0$  and kh=1. All panels show variation as a function of propagation direction  $\theta$  for three different values of  $\kappa h$ . (a) Comparison of (18) and the 3DKC, (22). (b) Comparison of (18) and the naïve estimate  $\tilde{c}_s$ . (c) Parameters  $\delta$  and  $\Delta$  as functions of  $\theta$ .

shown in Figure 4c. At the angles where the discrepancy is greatest ( $\theta \approx 0.2\pi$ ),  $\delta$  is large while  $\Delta$  is too small for a cancellation to occur according to (25). Our analysis strongly indicates that our model (18) should be superior here. In contrast, near  $\theta = \pi$  the difference between the two estimates is smaller than might be expected from the high  $\delta$  value, as is explained by the partial cancellation between  $\delta^2$  and  $\Delta$  in err<sub>KC</sub> in equation (28).

Finally Figure 4b shows clearly that the naïve estimate for  $\tilde{c}$  by extending the surface shear linearly into the deep using  $\tilde{c}_s$  (equation (20)) will work poorly in this case, where inclusion of the change of direction with depth is important for accurate dispersion estimation.

#### 4.4. Strongly Sheared Flows

We finally consider three strongly sheared examples where  $\delta$  is so large that the 3DKC fails for a range of wavelengths. Due to finite depth, phase velocities are limited in the long wave limit. Our first two profiles are of finite depth, and their variation over the water column exceeds the fastest possible wave speed, hence waves can never be considered "fast" nor the shear weak. The third is a deep water example with strong surface shear. All currents are unidirectional and we choose **k** along the current, an effectively 2D flow which is adequate as a demonstration. For all three example flows we have calculated  $\tilde{c}$  to an accuracy

better than  $10^{-4}$  using the piecewise linear approximation (Smeltzer & Ellingsen, 2017), allowing us to compare approximations to the  $\tilde{c}(k)$  which at the relevant level of accuracy can be considered "exact." In our examples we consider waves traveling downstream.

The first profile we consider is an example where our approximation (18) is successful whereas the 1st order 3DKC (22) fails completely:

$$U_1(z) = 3\sqrt{gh} \exp(z/h). \tag{33}$$

This profile is considered in panels a,d and g of Figure 5.  $U_1(z)$  has strong shear, and its variation  $|\Delta U_1|$  exceeds  $\tilde{c}$  for a large part of the wave number spectrum. The velocity variation over the column is not unreasonable for a region of rapid flow, for example over a local shallow in a river — a variation  $U_0 - U(-h) \sim 2\sqrt{gh}$  as here is obtained, for instance, for flow of surface velocity 4 m/s in 40 cm depth — or it might describe a fast film flow of depth 1cm with surface velocity 60 cm/s.

Although the 3DKC fails because  $\delta$  is too large, the 'near-potentiality' criterion (24) is very well satisfied, and  $c_{\approx}$  is within a few % of the correct value. Figures 5a, 5d, and 5g demonstrate the improved robustness of our approximation (18) compared to the Skop/Kirby and Chen model. The 3DKC becomes negative for long wavelengths, which is to say phase velocity should change sign in a system following the surface, which is physically unacceptable: the dispersion relation for a linear shear current always has one upstream and one downstream solution however strong the shear. For any smooth, nonlinear velocity profile, there always exists an even more strongly sheared linearly varying current, and since  $\tilde{c}_s$  from (20) is always positive, so must  $\tilde{c}$ .

Insight into the nature of second order correction terms is also offered by Figure 5d. The second-order correction to  $c_{\approx}$ , equation (19), is accurate to 1% or better for all k, while the second-order term as calculated from the expressions of Kirby and Chen (1989),  $\tilde{c}_{\text{KC},2nd}$ , also rectifies the deficiencies of its 1st order companion (the expression for  $\tilde{c}_{\text{KC},2nd}$  is somewhat bulky so is not quoted here) and works well; it has much higher relative error than (19), but this in itself is unlikely to be of practical importance. We are at present unable to explain the success of the second-order Kirby and Chen approximation in this case, and thus cannot determine under what conditions this holds true more generally. The criterion for the second-order approximation (19), on the other hand, is well controlled by condition (24).

We next consider a particularly difficult case where both shear and curvature are very large,

$$U_2(z) = \sqrt{ghexp(10z/h)}.$$
(34)

 $|\Delta U_2|$  exceeds  $\tilde{c}$  for much of the water column.  $U_2(z)$  is considered in Figure 5b, 5e, and 5h. Such a strongly sheared flow could occur locally, and might represent a realistic surface jet due to discharge of a fast flow into still waters, for example a jet speed of about 3m/s on a 1m depth.

We notice that for  $U_2(z)$  the parameter  $\Delta$  is about 0.22 for long waves, enough for the flow not to satisfy the 'near-potentiality' criterion (24). No approximation scheme based on near-potentiality can expect to fully succeed in this case. Since  $\delta$  exceeds unity, the 3DKC once again fails. Also  $c_{\approx}$  from (18) is inaccurate in this particular case, overestimating  $\tilde{c}$  by about 60% for long waves, yet the result is at least physically meaningful since the correct sign of  $\tilde{c}$  is guaranteed. For reference,  $\tilde{c}$  tends to the maximum  $\approx 0.5\sqrt{kh}$ as  $k \rightarrow 0$ . The second-order corrected approximation (19) moreover improves the estimates to the 20% level or better, whereas the second-order approximation due to Kirby and Chen (1989) is too small by about 40% for long waves and fares poorly for medium waves with  $kh \sim 5$  where  $\tilde{c}_{KC,2nd}$  is only about 20% of the real  $\tilde{c}$ . This particularly difficult example thus demonstrates the improved robustness of (18) and (19).

We finally consider another difficult case for any 'near-potentiality' approximation scheme, with strong surface shear in deep water,

$$U_{3}(z) = 3\sqrt{g/\alpha} \exp(\alpha z).$$
(35)

Since the shear is found near the surface only, the downstream phase velocity is unbounded for long waves, unlike for a linearly varying deep water current. Like in the previous cases, a region of wavelengths exists where the 3DKC yields an unphysical negative prediction,  $k/\alpha$  between approximately 0.2 and 1. The



**Figure 5.** Comparison of different approximation models for three strongly sheared velocity profiles. Pertain to (a, d, g)  $U_1(z)$  from equation (33), (b, e,h)  $U_2(z)$  from equation (34), and (c, f, i)  $U_3(z)$  from (35). (a-c) Velocity profiles. (d-f): 1st and second-order estimates using the present model to first and second order, respectively EL<sub>1st</sub> (equation (18)), and EL<sub>2nd</sub> (equation (19)), as well as the 1st and second-order approximations of Kirby and Chen (1989) (3DKC<sub>1st</sub> and 3DKC<sub>2nd</sub>, respectively), relative to the high accuracy calculation with the piecewise linear approximation (PLA). The legend in Figure 5d applies also to Figures 5e and 5f. (g, h, i): applicability parameters for the models as discussed in section 2.4.

robustness of our new approximation is thus demonstrated once again. It should be noted that while the 3DKC fails for medium wavelengths, it fares better than  $c_{\approx}$  and even  $c_{\approx,2nd}$  for  $k/\alpha \leq 0.05$ , despite the fact that  $\delta$  is as high as 0.6. Again this is because  $\Delta/\delta \approx \delta$  here, making criterion (25) well satisfied nevertheless, as seen in Figure 5i. This seems like a lucky coincidence, but given similar observations for the wind-driven profiles in section (4.1), it seems indicated that to some extent it holds more generally for a class of shear currents, a question which should be looked into in the future.

#### 5. Conclusions

We have presented a comprehensive theory of approximate dispersion relations for linear waves propagating atop a sub-surface horizontal shear current whose magnitude and direction can vary arbitrarily as a function of depth. We present a new analytical approximation which is shown to be more robust than the (3-D generalization of the) famous and widely used approximation by Skop (1987) and Kirby and Chen (1989), the 3DKC. The conditions of applicability of our new model are shown to be less restrictive, making it accurate in several realistic situation where the 3DKC is inaccurate or even breaks down. For the cases when the criteria for the 3DKC to work are satisfied, the two models are essentially equally good. Our approximations are tested for a range of different example flows.

A second-order accurate expression in terms of the suitable small parameter is also derived, and shown to greatly improve accuracy. The inclusion of the next order may not be called for in many situations, but constitutes a significant improvement in the more difficult situations discussed. The next order correction is robust and has the same condition of applicability as the first order (Kirby & Chen (1989) also derive a second-order correction to their model, whose conditions of applicability are not known at present).

Our thorough perturbation analysis is able to explain, for the first time, the success of the 3DKC for many cases, including situations where the assumptions behind its original derivation are strongly violated. Indeed, careful criteria are derived for our approximation, as well as for that of Kirby and Chen, to be accurate.

To leading order our new approximation involves essentially identical computational effort and complexity as that of Kirby and Chen. Our experience, however, is that our new second-order correction is significantly less complicated to implement than its sibling derived by said authors (while admitting that this could be a point of preference), and arguably more physically transparent.

#### **Appendix A: A Further Discussion and Mathematical Details**

#### **A1. Infinite Depth Expressions**

We here list key quantities in the limit  $kh \to \infty$ . Now  $c_0 = \sqrt{q/k}$ , and

$$\delta = \int_{-\infty}^{0} \frac{\mathbf{k} \cdot \mathbf{U}'(z)}{kc_0} e^{2kz} dz; \quad \tilde{U} = 2 \int_{-\infty}^{0} \mathbf{k} \cdot \mathbf{U}(z) e^{2kz} dz.$$
(A1)

The relation  $k\tilde{U} = \mathbf{k} \cdot \mathbf{U}_0 - kc_0 \delta$  still holds. The expression for  $\Delta(\tilde{c})$  simplifies greatly, to

$$\Delta(\tilde{c}) = -\frac{\tilde{c}}{kc_0^2} \int_{-\infty}^0 \frac{\mathbf{k} \cdot \mathbf{U}''(z)}{\mathbf{k} \cdot \Delta \mathbf{U}(z) - k\tilde{c}} [\tilde{U} - \tilde{u}(z)] e^{2kz} dz;$$
(A2)

$$\tilde{u}(z) = 2 \int_{-\infty}^{z} \mathbf{k} \cdot \mathbf{U}(\zeta) e^{2k(\zeta - z)} d\zeta.$$
(A3)

The first two solutions to w(z) are

$$w^{(0)}(z) = w^{(0)}(0)e^{kz}; \ w^{(1)}(z) = \frac{w^{(0)}(0)}{k} \int_{-\infty}^{z} \frac{\mathbf{k} \cdot \mathbf{U}'(\zeta)}{\mathbf{k} \cdot \Delta \mathbf{U}(\zeta) - k\tilde{c}} e^{k\zeta} \sinh k(z-\zeta) d\zeta.$$
(A4)

For the purposes of the approximations for group velocity,

$$C_{L} = \int_{-\infty}^{0} \mathbf{k} \cdot \mathbf{U}'(z) (1 + 2kz) e^{2kz} / k.$$
 (A5)

#### A2. Simplification of $\Delta(\tilde{c})$

Here follow the details of the simplification of  $\Delta(\tilde{c}) = \epsilon \Omega_l + \epsilon \Omega_{k1} \delta + \epsilon \Omega_{k1}$  in equations (10) and (11) to the form (14). Using  $\delta = (\mathbf{k} \cdot \mathbf{U}_0 - k\tilde{U})/kc_0$  from (22) and sinh  $kh = 2\cosh kh/\sinh 2kh$ , we write

$$\delta \varepsilon \Omega_{K1} = \frac{2\tilde{c}}{k^2 c_0^2} \int_{-h}^{0} \Upsilon(z) \frac{\sinh k(z+h)}{\sinh 2kh} 2\sinh kz \cosh kh [\mathbf{k} \cdot \mathbf{U}_0 - k\tilde{U}] dz$$
(A6)

where the shorthand  $\Upsilon(z) = \mathbf{k} \cdot \mathbf{U}''(z)/(\mathbf{k} \cdot \Delta \mathbf{U} - k\tilde{c})$  is introduced. Next we perform a partial integration of the inner integral in  $\epsilon \Omega_{K2}$  to obtain

$$\varepsilon \Omega_{K2} = \frac{2\hat{c}}{k^2 c_0^2} \int_{-h}^{0} \Upsilon(z) \frac{\sinh k(z+h)}{\sinh 2kh} [\mathbf{k} \cdot \mathbf{U}_0 \sinh k(h-z) - \mathbf{k} \cdot \mathbf{U}(z) \sinh k(h+z) - 2k \int_{z}^{0} \mathbf{k} \cdot \mathbf{U}(\zeta) \cosh k(2\zeta+h-z) d\zeta] dz.$$
(A7)

Using that  $\sinh k(h-z) - \sinh k(h+z) + 2\sinh kz \cosh kh = 0$ , we find

$$\Delta = -\frac{2\tilde{c}}{kc_0^2} \int_{-h}^{0} \Upsilon(z) \frac{2\sinh k(z+h)}{\sinh 2kh} [\tilde{U}\sinh kz\cosh kh + \int_{z}^{0} \mathbf{k} \cdot \mathbf{U}(\zeta)\cosh k(2\zeta+h-z)d\zeta]dz,$$
(A8)

which takes the form (14) with minimal further manipulation.

#### A3. The Short-Wave Approximation of Shrira

We consider the approximation due to Shrira (1993) derived for short waves. For large *k* the integral is dominated by  $|z| \le 1/2k$ , and we now assume the wave short enough that a near–surface Taylor expansion  $\mathbf{U}(z) \approx \mathbf{U}_0 + \mathbf{U}'_0 z + \cdots$  is in order. Terms proportional to  $\mathbf{U}''(z)$  are then small, and the integral term *I* (see equation (4)) is a small correction in equation (3). Since  $\Delta \mathbf{U} \approx \mathbf{U}'_0 z$  we have  $\mathbf{k} \cdot \Delta \mathbf{U} \le \mathbf{U}'_0/2k$  which is negligible compared to  $k\tilde{c} \sim \sqrt{gk}$ . We thus obtain  $I \approx -(2c_0/\tilde{c})\delta_5$  with the short-wave smallness parameter

$$\delta_{S} = \int_{-h}^{0} \frac{\mathbf{k} \cdot \mathbf{U}''(z) \sinh^{2}k(z+h)}{k^{2}c_{0}\sinh 2kh} dz$$
(A9)

which is a depth-averaged velocity profile curvature. Equation (3) becomes approximately

$$\tilde{c}^2 - 2\tilde{c}c_0\delta_S + k^{-2}\tilde{c}\mathbf{k}\cdot\mathbf{U}_0$$
tanh  $kh - c_0^2 = 0$ .

Assuming  $\delta_s$  to be small and expanding to first order gives

$$\tilde{\epsilon} \approx c_{\rm s}(1+c_0\delta_{\rm S}/\sqrt{\cdots})$$
 (A10)

where  $\sqrt{\cdots}$  is the square root term in (20). This is the leading order of the approximation of Shrira (1993).

The criterion of applicability,  $\delta_S \ll 1$ , only in general holds for large *k*. For longer waves a cancellation occurs in (3), because, by partial integration,

$$\delta_{\rm S} = \mathbf{k} \cdot \mathbf{U}_0' \tanh kh/(2k^2c_0) - \delta.$$

For short waves both terms on the right hand side are small. For longer waves, the first term no longer is, but it is cancelled by another term in (3). If  $\delta$  is also small for long waves, this cancellation effectively replaces the small parameter  $\delta_s$  for short waves, by  $\delta$  for longer waves. Approximation (A10) has no such cancellation, and therefore fails for longer waves, while approximations (18) and (22) typically do not.

#### A4. Stability of a Critical Layer

Let there be a critical layer at  $z = z_{cr}$  so that near this depth,  $\mathbf{k} \cdot \mathbf{U}(z) \approx kc + (z-z_c)\mathbf{k} \cdot \mathbf{U}'_{cr}$  where  $\mathbf{U}_c = \mathbf{U}(z_c) = c$ and  $\mathbf{U}'_c = \mathbf{U}'(z_c)$ . The integral in (3) now has a pole on the axis of integration. It becomes well defined when considered as the  $t \to \infty$  limit of a corresponding initial value problem (Peregrine, 1976). Suppose the wave has been made by a wave paddle with frequency  $\omega = kc$ , and whose amplitude has increased slowly from zero at  $t = -\infty$ . The generated wave will have time dependence  $\exp(-i\omega t + \epsilon t)$  with  $\epsilon = 0^+$ . The effect is to replace  $kc \to kc + i\epsilon$ , moving the pole slightly off the *z* axis. Approximating  $w \approx w^{(0)}$  from (9a) and using the Sokhotski-Plemelj theorem, a 'near–potentiality' estimate of *l* from (4) is

$$I \approx \mathcal{P} \int_{-h}^{0} \frac{2\mathbf{k} \cdot \mathbf{U}'' \sinh^2 k(z+h) dz}{k(\mathbf{k} \cdot \Delta \mathbf{U} - k\tilde{c}) \sinh 2kh} + \frac{2\pi i \mathbf{k} \cdot \mathbf{U}'_c \sinh^2 k(z_c+h)}{k |\mathbf{k} \cdot \mathbf{U}_c| \sinh 2kh} \equiv I_r + iI_i.$$
(A11)

 $\mathcal{P}$  denotes the principal value. With this, the first-order version of (13) becomes  $\tilde{c}^2 + 2c_0\tilde{c}\delta - c_0^2 + i\tilde{c}^2I_i = 0$ . Now suppose the complex  $\tilde{c}$  is  $\tilde{c} = \tilde{c}_r + ic_i$ , and noting that  $c_i$  is order  $I_i$ , we solve for the real and imaginary parts of the resulting equation to order  $I_i$ . This results in  $\tilde{c}_r \approx \tilde{c}$ , and

$$c_i \approx -\tilde{c}_{\approx}^2 I_i / \left( 2c_0 \sqrt{1+\delta^2} \right). \tag{A12}$$

Here,  $\tilde{c}_r$  is approximated using, e.g., (18). For the case of short waves in deep water this agrees with the result of Shrira (1993). Now,  $c_i > 0$  implies unstable flow. Since  $c_0 > 0$ , instability is predicted when  $\mathbf{k} \cdot \mathbf{U}'_c < 0$ .

#### References

Abdullah, A. J. (1949). Wave motion at the surface of a current which has an exponential distribution of vorticity. The Annals of the New York Academy of Sciences, 51, 425–441.

Banihashemi, S., Kirby, J. T., & Dong, Z. (2017). Approximation of wave action flux velocity in strongly sheared mean flows. Ocean Modelling, 116, 33–47.

Bender, C. M., & Orszag, S. A. (1978). Advanced mathematical methods for scientists and engineers. New York, NY: McGraw-Hill.

Burns, J. C. (1953). Long waves in running water. Proceedings of the Cambridge Philosophical Society, 49, 695–706.

Campana, J., Terrill, E. J., & de Paolo, T. (2017). A new inversion method to obtain upper–ocean current–depth profiles using X-band observations of deep–water waves. *Journal of Atmospheric and Oceanic Technology*, 34, 957–970.

Craik, A. D. D. (1968). Resonant gravity-wave interactions in a shear flow. Journal of Fluid Mechanics, 34, 531-549.

Dalrymple, R. A. (1977). A numerical model for periodic finite amplitude waves on a rotational fluid. *Journal of Computational Physics*, *24*, 29–42. Dong, Z., & Kirby, J. T. (2012). Theoretical and numerical study of wave-current interaction in strongly-sheared flows. In P. Lynett & J.M. Smith (Eds.), *Proceedings of 32nd Conference on Coastal Engineering*. Santander, Spain.

Drazin, P. G., & Reid, W. H. (1981). Hydrodynamic stability. Cambridge, UK: Cambridge University Press.

Elias, E. P. L., Gelfenbaum, G., & Van der Westhuysen, A. J. (2012). Validation of a coupled wave-flow model in a high-energy setting: The mouth of the Columbia river. *Journal of Geophysical Research*, *117*, C09011. https://doi.org/10.1029/2012JC008105

Ellingsen, S. Å. (2014). Initial surface disturbance on a shear current: The Cauchy–Poisson problem with a twist. *Physics of Fluids*, 26, 082104.
Ellingsen, S. Å. (2016). Oblique waves on a vertically sheared current are rotational. *European Journal of Mechanics B: Fluids*, 56, 156–160.
Ellingsen, S. Å., & Brevik, I. (2014). How linear surface waves are affected by a current with constant vorticity. *European Journal of Physics*, 35(2), 025005.
Fenton, J. D. (1973). Some results for surface gravity waves on shear flows. *Instituteof Mathematics and its Applications*, 12, 1–20.
Hunt, J. N. (1955). Gravity waves in flowing water. *Proceedings of the Royal Society of London A: Mathematical, Physical, and Engineering Sciences*, 231, 493–504.

Karageorgis, P. (2012). Dispersion relation for water waves with non-constant vorticity. European Journal of Mechanics B/Fluids, 34, 7–12. Kirby, J. T., & Chen, T.-M. (1989). Surface waves on vertically sheared flows: Approximate dispersion relations. Journal of Geophysical

Research, 94, 1013–1027.

Ko, J., & Strauss, W. (2008a). Effect of vorticity on steady water waves. Journal of Fluid Mechanics, 608, 197–215.

Ko, J., & Strauss, W. (2008b). Large–amplitude steady rotational water waves. European Journal of Mechanics B/Fluids, 27, 96–109.
Kumar, N., Voulgaris, G., & Warner, J. C. (2011). Implementation and modification of a three–dimensional radiation stress formulation for surf zone and rip–current applications. Coastal Engineering, 58, 1097–1117.

Kumar, N., Voulgaris, G., Warner, J. C., & Olabarrieta, M. (2012). Implementation of the vortex force formalism in the coupled ocean-atmosphere-wave-sediment transport (COAWST) modeling system for inner shelf and surf zone applications. Ocean Modelling, 47, 65–95.

Lund, B., Graber, H. C., Tamura, H., Collins, C. O., III, & Varlamov, S. M. (2015). A new technique for the retrieaval of near-surface vertical current shear from marine X-band radar images. *Journal of Geophysical Research Oceans, 120*, 8466–8486. https://doi.org/10.1002/2015JC010961

Moreira, R. M., & Chacaltana, J. T. A. (2015). Vorticity effects on nonlinear wave-current interactions in deep water. Journal of Fluid Mechanics, 778, 314–334.

Morland, L. C., Saffman, P. G., & Yuen, H. C. (1991). Waves generated by shear layer instabilities. Proceedings of Royal Society of London A: Mathematical, Physical, and Engineering Sciences, 433, 441–450.

Nwogu, O. G. (2009). Interaction of finite-amplitude waves with vertically sheared current fields. *Journal of Fluid Mechanics*, 627, 179–213. Peregrine, D. H. (1976). Interaction of water waves and currents. *Advances in Applied Mechanics*, 16, 9–117.

Quinn, B. E., Toledo, Y., & Shrira, V. I. (2017). Explicit wave action conservation for water waves on vertically sheared flows. Ocean Modelling, 112, 33–47.

Rayleigh, L. (1892). On the question of the stability of the flow of fluids. Philosophical Magazine, 34, 59–70.

Shrira, V. I. (1993). Surface waves on shear currents: Solution of the boundary-value problem. Journal of Fluid Mechanics, 252, 565–584.

Simmen, J. A., & Saffman, P. G. (1985). Steady deep-water waves on a linear shear current. *Studies in Applied Mathematics*, 73, 35–57. Skop, R. A. (1987). Approximate dispersion relation for wave-current interactions. *Journal of Waterway, Port, Coastal, and Ocean Engineering*, 113, 187–195.

Smeltzer, B. K., & Ellingsen, S. Å. (2017). Surface waves on arbitrary vertically-sheared currents. *Physics of Fluids*, *29*, 047102.
 Stewart, R. J., & Joy, J. W. (1974). HF radio measurements of surface currents. *Deep Sea Research and Oceanographic Abstracts*, *21*, 1039–1049.
 Swan, C., & James, R. L. (2000). A simple analytical model for surface water waves on depth-varying current. *Applied Ocean Research*, *22*, 331–347.
 Taylor, G. (1955). The action of a surface current used as a breakwater. *Proceedings of the Royal Society London A: Mathematical, Physical, and Engineering Sciences*, *231*, 466–478.

Teles da Silva, A., & Peregrine, D. H. (1988). Steep, steady surface waves on water of finite depth with constant vorticity. *Journal of Fluid Mechanics*, 195, 281–302.

Thompson, P. D. (1949). The propagation of small surface disturbances through rotational flow. Annals of the New York Academy of Sciences, 51, 462–474.

Velthuizen, H. G. M., & van Wijngaarden, L. (1969). Gravity waves over a non-uniform flow. *Journal of Fluid Mechanics*, *39*, 817–829. Young, W. R., & Wolfe, C. L. (2014). Generation of surface waves by shear-flow instability. *Journal of Fluid Mechanics*, *739*, 276–307. Zhang, X. (2005). Short surface waves on surface shear. *Journal of Fluid Mechanics*, *541*, 345–370.

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